

# **Evaluation of the Multi-Angle Implementation of Atmospheric Correction (MAIAC) Aerosol Algorithm through Intercomparison with VIIRS Aerosol Products and AERONET**

Stephen D. Superczynski

Systems Research Group Inc., NOAA/NESDIS/STAR, College Park, Maryland

Shobha Kondragunta

NOAA/NESDIS/STAR, College Park, Maryland

Alexei I. Lyapustin

NASA Goddard Space Flight Center, Greenbelt, Maryland

Corresponding author: Stephen Superczynski, SRG Inc. (at NOAA/NESDIS/STAR),

5825 University Research Court, College Park, MD 20740

E-mail: [stephen.superczynski@noaa.gov](mailto:stephen.superczynski@noaa.gov)

## **Key Points:**

- MAIAC algorithm is evaluated for use in future satellite missions to derive information on aerosol properties
- MAIAC's use of time-series observations allow it to derive BRF which in turn improves cloud masking and aerosol-surface retrievals.
- Comparison with AOT from VIIRS and AERONET show that MAIAC exhibits low bias over North America with high spatial coverage

25    **Abstract**

26    The Multi-Angle Implementation of Atmospheric Correction (MAIAC) algorithm is under evaluation  
27    for use in conjunction with the Geostationary Coastal and Air Pollution Events (GEO-CAPE) mission.  
28    Column aerosol optical thickness (AOT) data from MAIAC are compared against corresponding data  
29    from the Visible Infrared Imaging Radiometer Suite (VIIRS) instrument over North America during  
30    2013. Product coverage and retrieval strategy, along with regional variations in AOT through  
31    comparison of both matched and un-matched seasonally gridded data are reviewed. MAIAC shows  
32    extended coverage over parts of the continent when compared to VIIRS, owing to its pixel selection  
33    process and ability to retrieve aerosol information over brighter surfaces. To estimate data accuracy,  
34    both products are compared with AERONET Level 2 measurements to determine the amount of error  
35    present and discover if there is any dependency on viewing geometry and/or surface characteristics.  
36    Results suggest that MAIAC performs well over this region with a relatively small bias of -0.01;  
37    however there is a tendency for greater negative biases over bright surfaces and at larger scattering  
38    angles. Additional analysis over an expanded area and longer time period are likely needed to determine  
39    a comprehensive assessment of the products capability over the Western Hemisphere.

40    **Index Terms:**

41            Aerosols and particles

42            Remote sensing

43    **Keywords:**

44            MAIAC, Suomi-NPP VIIRS, Aerosol Optical Thickness, Intercomparison, Evaluation

45

46

## 47    **1. Introduction**

48    Aerosols are a key component of the Earth's climate and environmental system due to their impact on  
49    the radiative budget of the planet and influence on air quality events [*Ramanathan et al.*, 2001].  
50    Information on the amount and composition of the aerosol particles suspended in the atmosphere is  
51    required to understand their role as both direct contaminants and precursors to air pollution [*Wang and*  
52    *Christopher*, 2003; *Al-Saadi et al.*, 2005]. The GEO-CAPE mission was recommended by the National  
53    Research Council's 2007 Decadal Survey in order to provide multiple observations per day in support of  
54    the atmospheric composition and coastal biophysics disciplines [*NRC*, 2007]. Many current sensors  
55    dedicated toward atmospheric composition sit in Low Earth Orbit (LEO) and have only one daytime and  
56    one nighttime overpass for a given location when more frequent measurements are needed to fully  
57    monitor the emission of pollutants and their transport. A geostationary platform provides both the  
58    temporal and spatial resolution needed to understand the conditions and processes leading to poor air  
59    quality events and the necessary response [*Lahotz et al.*, 2012].

60    Originally planned as a large satellite carrying multiple instruments, GEO-CAPE has shifted toward a  
61    phased implementation making use of available space on commercial geostationary satellites. This  
62    utilization of hosted payloads should help to reduce risk and costs, and has been supported by both  
63    science working groups [*Fishman et al.*, 2012]. The atmospheric science working group is tasked with  
64    developing a strategy which allows for the observation of aerosols and trace gases for use in air quality  
65    studies. The MAIAC algorithm is the current candidate to provide information on aerosols from this  
66    geostationary satellite.

67    The MAIAC algorithm provides simultaneous retrievals of surface bidirectional reflectance distribution  
68    function (BRDF), bidirectional reflectance factor (BRF) commonly called surface reflectance, and AOT  
69    at 466 nm over clear sky and snow-free scenes using a time series of MODerate Imaging

70 Spectroradiometer (MODIS) observations. This BRDF characterization over time for varying  
71 geometries is used, along with the spectral regression coefficient (SRC), to help the MAIAC algorithm  
72 retrieve AOT over bright surfaces with improved accuracy [Lyapustin *et al.*, 2011].

73 Here this new generic algorithm is assessed through a comparison with the operational VIIRS aerosol  
74 algorithm which uses an atmospheric correction approach. VIIRS was chosen for this comparison due to  
75 the improvements over its predecessors in terms of resolution, pixel aggregation, and swath width. For  
76 instance, MODIS has a long history of providing aerosol retrievals with high accuracy, but it currently  
77 only produces AOT at a maximum resolution of 3 km, and has greater distortion at the swath edge when  
78 compared to VIIRS. The Multi-angle Imaging Spectroradiometer (MISR) uses nine fixed-angle cameras  
79 to view each location at a variety of viewing angle which allows it to also retrieve AOT over brighter  
80 surfaces; however its limited swath width (400 km) and coarse resolution (17.6 km) are prohibitive to its  
81 inclusion in this analysis. Ultimately, the sensor characteristics and availability of .75 km AOT retrievals  
82 make it ideal for a comparison with MAIAC. In this study, a years' worth of AOT from both MAIAC  
83 and VIIRS over the North American continent is analyzed to look at differences in cloud screening, bias  
84 dependence and overall accuracy.

## 85 **2. Data**

### 86 **2.1 MAIAC AOT**

87 The MAIAC algorithm retrieves surface reflectance and AOT using MODIS L1B reflectances which  
88 have been gridded at a 1 km resolution. It utilizes a 4-16 day time series of clear MODIS scenes to  
89 retrieve BRDF and Spectral Regression Coefficients (SRC), which relates surface reflectance at  
90  $0.466\mu\text{m}$  and  $2.13\mu\text{m}$  (MODIS bands 3 and 7) [Lyapustin *et al.*, 2012]. Unlike MISR, which collects  
91 nearly-simultaneous observations of each pixel from various angles, the MAIAC algorithm uses  
92 consecutive overpasses from a single-look instrument like MODIS to acquire multi-angle sets of

93 observations for each location. The use of a time-series of gridded MODIS observations also has the  
94 advantage of being able to simulate geostationary satellite observations, albeit with a significantly larger  
95 time difference between images. MAIAC relies on the assumption that surface reflectance changes  
96 rapidly in space but slowly in time, and therefore can be assumed constant over limited time scales. By  
97 contrast, the extent of clouds and aerosols can change greatly between MODIS overpasses.

98 The following is a brief overview of the MAIAC aerosol algorithm, a more detailed description of the  
99 MAIAC theoretical background and processing steps can be found in *Lyapustin et al.*, (2011). Once the  
100 MODIS reflectance is gridded and split into both 600 x 600 km tiles and 25 x 25 km blocks, they are  
101 placed in a queue of 4-16 days. Water vapor is first derived from MODIS near-IR bands [*Lyapustin et al.*,  
102 2014] using a modification of the algorithm described in *Gao and Kauffman* (2003). An internal cloud  
103 mask uses spectral reflectance and brightness temperature tests similar to the operational MODIS cloud  
104 mask algorithm [*Frey et al.*, 2008], along with the reference clear-sky image developed using a covariance  
105 based algorithm. Clouds can be detected since the spatial pattern of the surface often doesn't change  
106 noticeably from day to day, while cloud residency is relatively short. Scenes are compared at both the  
107 block and pixel level against a clear-sky reference image built using the data queue [*Lyapustin et al.*,  
108 2008]. The BRDF is then retrieved at MODIS band 7 (2.1  $\mu\text{m}$ ) for clear pixels, followed by retrieval of  
109 SRC in MODIS band 3 (0.466 $\mu\text{m}$ ). This retrieval of SRC gives an assessment of surface BRDF (0.466 $\mu\text{m}$ )  
110 at pixel level, which allows MAIAC to retrieve AOT at high 1km resolution.

111 The MAIAC algorithm provides AOT at 466 nm, however in order to compare directly with VIIRS, it  
112 must be converted to AOT at 550 nm. To do this, a set of ratios representing the spectral slope of a given  
113 AOT are used. These ratios, which are taken directly from the aerosol background model, are part of the  
114 MAIAC look-up tables [*Lyapustin et al.*, 2011]. MODIS-based MAIAC aerosol products were produced  
115 over North America for the entire MODIS record up until July 2014. MAIAC is currently at version 1,  
116 and data used for this analysis was obtained from NASA on November 17, 2014.

## 117    **2.2 VIIRS AOT**

118    The Visible and Infrared Imaging Radiometer Suite (VIIRS) is a scanning radiometer carried on board  
119    the Suomi-NPP (National Polar-orbiting Partnership) satellite; a joint venture between NOAA and  
120    NASA meant to help transition to the Joint Polar Satellite System (JPSS), the next generation in U.S.  
121    polar-orbiting satellites. The operational VIIRS AOT product is produced by the Interface Data  
122    Processing System (IDPS), which takes raw instrument data from S-NPP and processes them into the  
123    Sensor Data Records (SDRs) that are used as inputs for the Environmental Data Records (EDRs),  
124    including AOT. The aerosol algorithm uses the dark-target approach to retrieve AOT. This method is  
125    built upon the legacy of retrieving aerosol properties from previous earth sensing satellite missions  
126    [Holben *et al.*, 1992; Kaufman *et al.*, 1997]. The algorithm is comprised of two distinct parts which are  
127    applied based on the surface type. Over ocean, the VIIRS algorithm is nearly identical to the MODIS  
128    ocean algorithm [Tanre *et al.*, 1997], which uses a combination of fine and coarse mode aerosol models  
129    in attempt to replicate the top-of-atmosphere (TOA) reflectance. Over land, the VIIRS aerosol  
130    algorithm is based on the MODIS Atmospheric Correction algorithm for determining surface reflectance  
131    [Vermote and Kotchenova, 2008]. Aerosol information is retrieved by comparing the derived spectral  
132    surface reflectance ratios to prescribed ratios of those reflectances, and chooses the aerosol model and  
133    AOT that minimizes the residual. The VIIRS aerosol algorithm operates under the assumption of a  
134    Lambertian surface when retrieving the surface reflectance. An overview of the VIIRS sensor and an in-  
135    depth explanation of the scientific background and flow of the VIIRS aerosol algorithm are presented in  
136    Jackson *et al.*, (2013).

137    The aerosol retrieval for both ocean and land is performed at the pixel resolution (750 m). This pixel  
138    level product is known as the Intermediate Product (IP) as it is used to create the aggregated AOT EDR,  
139    along with acting as an input for other VIIRS products. The VIIRS algorithm aggregates 8x8 arrays of  
140    IP AOT pixels into a single EDR pixel with a resolution of 6 km at nadir. At the IP level, the VIIRS

141 Cloud Mask (VCM) and a series of internal checks are applied to the aerosol product, resulting in each  
142 pixel being given one of four quality designations. AOT is reported only for pixels in the two best  
143 quality levels (good and degraded) and therefore these are the only pixels included in the aggregation  
144 process, which also incorporates additional filtering and internal checks, producing a higher quality  
145 product.

146 A full year of VIIRS IP AOT spanning the time from February 1, 2013 to February 1, 2014 was used to  
147 compare against the MAIAC product. The selection of this time period was predicated by data  
148 availability and maturity. The VIIRS Aerosol algorithm has undergone multiple upgrades since launch  
149 to improve the accuracy and precision of its retrievals. One significant upgrade was a change to the  
150 spectral reflectance ratios used in the land inversion which took place in January 2013 [*Hongqing et al.,*  
151 *2013*]. This greatly reduced the bias in the aerosol products over land and allowed the product to reach  
152 ‘validated’ status. Because data prior to this change becoming operational are still considered  
153 ‘provisional’, they were not included in this analysis. Officially, the version of the product used in this  
154 study was given a maturity level of Validated Stage II in August 2014, meaning that it has been shown  
155 to meet the performance thresholds [*NOAA-NESDIS, 2014*] using a moderate set of test data. There are  
156 no such standards for the IP product; however it also meets the EDR requirements, making it suitable for  
157 quantitative analysis.

158 Other significant changes have occurred to the AOT product after the time period used in this study  
159 which had impacts on retrieval accuracy and to a lesser extent, spatial coverage. These include an  
160 improvement in snow screening, spatial homogeneity tests, and the removal of the ephemeral water test  
161 which often incorrectly screened out portions of heavy smoke plumes. Unfortunately due to the MAIAC  
162 data record ending in mid-2014, data containing these fixes were not included in this analysis.

## 163 **2.3 AERONET**

164 AERONET is a global network of ground-based, automatic sky-scanning spectral radiometers used to  
165 measure aerosol optical properties [Holben *et al.*, 1998]. Developed and maintained by NASA, these  
166 weather resistant sun photometers are a vital source of information for aerosol research and the  
167 validation of satellite derived aerosol properties. The direct-sun measurements are used to compute the  
168 column AOT at a variety of wavelengths from 340 – 1020 nm, spanning a majority of the visible and  
169 Near-IR spectrum. Angstrom Exponent (AE) is also retrieved using wavelength pairs in the  
170 aforementioned range, along with the column water vapor. Level 2.0 AOT from AERONET sites in  
171 North America are used to compare against both the MAIAC and VIIRS AOT to determine accuracy  
172 and uncover any bias dependencies. Level 2 data has the highest quality assurance of all AEROENT  
173 data and is cloud-cleared and fully calibrated [Smirnov *et al.*, 2000]. The “ground truth” AOT at the  
174 VIIRS and MAIAC wavelengths are computed using the AERONET AOT at 500 and 440 nm  
175 respectively, using the AE retrieved in the 440-675 nm range.

## 176 **2.4 CALIPSO**

177 The Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) is an active lidar instrument aboard  
178 the CALIPSO satellite. It provides vertically resolved information on clouds and aerosols using profiles  
179 of attenuated backscatter at 532 and 1064 nm at an along track resolution of 333 meters and a vertical  
180 resolution of 30 meters [Winker *et al.*, 2009]. CALIOP is able to detect the number and extent of  
181 features such as aerosol or cloud layers using the backscatter profiles [Vaughan *et al.*, 2004]. The level 2  
182 products are produced at the nominal resolution of 333 m as well as 1 and 5 km by aggregating  
183 consecutive observations. For this study, the 1 km cloud layer products are used to verify the accuracy  
184 of the MAIAC and VIIRS cloud masks and determine if any issues related to cloud screening are  
185 influencing the analysis. A binary cloud mask is constructed from the ‘Number of Layers Found’  
186 dataset, which simply gives the number of cloud layers found within that 1 km profile.



## 187    **3. Results and Discussion**

### 188    **3.1 Daily gridding of VIIRS and MAIAC**

189    Before assessing the MAIAC algorithm and how it compares to VIIRS, the datasets were gridded to  
190    directly compare their spatial extent and the quality of AOT retrievals. A grid was constructed with a  
191    0.25° resolution in order to capture as much of the AOT spatial variability while limiting computational  
192    cost. The shaded domain outlined in Figure 1 shows the extent of the grid whose domain is limited by  
193    the MAIAC coverage over North America, which is largely confined to the Continental U.S. and  
194    Mexico. The result is a grid with dimensions of 256 x 116, or a total of 29,696 grid boxes.

195    In order to compare the best retrievals from both algorithms, a set of quality checks were applied during  
196    the gridding process. To start, data from both algorithms are restricted to the highest quality retrievals  
197    over land. To avoid any possible cloud leakage, the candidate pixel was required to be confidently clear  
198    and not be adjacent to a cloudy pixel in order to be used for gridding. Both MAIAC and VIIRS AOT  
199    have an associated geolocation file which gives the center coordinates of each pixel. The gridding  
200    process averages any valid pixels whose center lat/lon falls within the same grid box, and the number of  
201    observations included in that average is recorded. These daily gridded datasets were then averaged to  
202    look at statistics on the monthly to seasonal scale.

### 203    **3.2 Direct Comparison**

204    Once gridding of the data was completed, the datasets were directly compared through analysis of un-  
205    paired seasonal AOT and looking at the differences in retrieval numbers. Due to the ability of MAIAC  
206    to retrieve AOT over brighter surfaces, it was expected that it would have greater spatial coverage than  
207    the operational VIIRS product, particularly in areas of sparse vegetation.

#### 208    **3.2.1 Data coverage**

209 Seasonal averages of AOT from MAIAC and VIIRS and the total number of retrievals per grid were  
210 analyzed in order to get a sense of the differences in coverage, and gain insight into the retrieval strategy  
211 and cloud screening of each algorithm. Figure 2 provides a look at the average of AOT (top) and  
212 number of retrievals per grid (bottom) per season for each dataset. MAIAC has greater coverage and  
213 more retrievals than VIIRS particularly across the western half of the CONUS. MAIAC coverage is  
214 nearly complete during the summer and fall seasons, save for some inland water bodies and regions such  
215 as Great Salt Flats (UT) and White Sands (NM), while VIIRS is not able to retrieve over the bright  
216 surfaces that make up a large portion of the western U.S. This disparity in coverage is seen across all  
217 seasons with the differences being greater during winter and spring due to seasonal phenology. There are  
218 some similarities however; for instance during winter when neither MAIAC or VIIRS retrieve enough to  
219 populate grids over the northernmost sections of the U.S. or the high altitude regions of the inter-  
220 mountain west. The reason for this is likely a combination of the solar zenith angle limits placed on  
221 good quality data and near-constant snow cover in these regions during the cold season.

222 In terms of actual AOT values, Figure 2c highlights some differences between MAIAC and VIIRS.  
223 While the spatial patterns are very similar between the two, VIIRS tends to retrieve slightly higher AOT  
224 over many regions. Over urban areas or mountainous terrain, this difference can be quite large and is  
225 noticeable in many seasons. In the springtime months, VIIRS AOT is also higher in the upper Mid-west  
226 and Great Lakes region where melting snow is likely contaminating the pixels leading to a poor  
227 retrieval. These anomalies associated with sub-pixel snow have since been addressed in the operational  
228 VIIRS algorithm.

229 Looking collectively at the results of this comparison, there are some features present in multiple  
230 seasons which emphasize the differences between the two algorithms and their pixel selection strategy.  
231 The underlying surface reflectance plays an important role in coverage of both datasets. MAIAC has  
232 shown the ability to retrieve AOT over the bright and soil dominated surfaces that are present across

233 much of the western U.S., while VIIRS is only able to retrieve over darker or vegetated regions. This is  
234 also a problem in regions with high agricultural activity, such as the Lower Mississippi River Basin  
235 where fallow land prevents VIIRS from consistently retrieving AOT in all seasons besides the primary  
236 growing season (JJA). However surface reflectance alone cannot account for the differences in  
237 retrievals seen in many other parts of the US throughout the year.

### 238 **3.2.2 Cloud Screening**

239 In an effort to understand the difference in coverage and to determine how the cloud masks are  
240 performing, data from MAIAC and VIIRS were collocated with the CALIOP instrument aboard the  
241 CALIPSO satellite. First, the two cloud masks are converted to a binary mask with either a ‘clear’ or  
242 ‘cloudy’ designation. All datasets are subsetting to regions of overlap, after which the closest  
243 MAIAC/VIIRS pixel to the CALIOP profile is found using a modified version of the nearest neighbor  
244 approach utilized in similar comparison studies [*Heidinger et al.*, 2012; *Kopp et al.*, 2014]. Here we use  
245 a time window of 10 minutes centered on the CALIOP observation time in order to avoid cases where  
246 clouds detected by CALIOP have moved out of the MAIAC/VIIRS field of view. A maximum allowed  
247 distance of one pixel width is used to ensure that the closest pixel is indeed chosen, this is particularly  
248 necessary where the CALIOP profile passes from one tile/granule to the next. Collocation results  
249 between the cloud masks and CALOP detection were compared and are presented in Table 1 as a  
250 confusion matrix.

251 Our first observation from Table 1 is that a considerably higher number of collocations for MAIAC exist  
252 than for VIIRS. This is not only due to MAIAC’s increased retrieval numbers but the use of reflectance  
253 data from MODIS, which is part of the A-train constellation [*Stephens et al.*, 2002] and shares a similar  
254 orbit and overpass time with CALIPSO. The VIIRS instrument flies at a slightly higher altitude and

255 therefore has a different orbital track, the consequence of which is a ground track that only coincides  
256 closely with the A-train satellites once every few days.

257 To help determine the performance of each set of matchups we look at overall accuracy (Equation 1)  
258 along with two additional statistical measures: the True Positive Rate (TPR), and True Negative Rate  
259 (TNR) for which the formulas are given in Equations 2 and 3, respectively. The abbreviations used in  
260 these equations are noted next to their respective statistics in Table 1. A high TPR value indicates that  
261 the cloud mask is able to limit the number of false negatives (type II error), which lead to cloud leakage  
262 in the resulting product. Conversely, TNR is a measure of how good the cloud mask is at reducing the  
263 number of false positives (type I error); these false alarms can reduce the number of high quality  
264 retrievals and introduce sampling biases.

$$265 \quad Accuracy = \frac{TP + TN}{TP + TN + FN + FP}$$

$$266 \quad TPR = \frac{TP}{TP + FN}$$

$$268 \quad TNR = \frac{TN}{TN + FP}$$

269  
270  
271 Overall accuracy of the both the MAIAC cloud mask (MCM) and the VCM were found to be identical  
272 (Table 1), but while the overall accuracy for the two cloud masks may be comparable, the errors  
273 observed were dissimilar. The TPR and TNR metrics highlight the different types of errors associated  
274 with each cloud mask. For instance, TPR for the MCM during this period is 96%, meaning that less  
275 than 5% of cloudy pixels were incorrectly designated as clear, while the TNR for MAIAC is only 72%,  
276 leaving over a quarter of the clear pixels as determined by CALIOP out of the AOT processing chain

277 due to the supposition they are cloudy. Monthly statistics for MAIAC show there is some seasonality to  
278 the TNR since it does not fall below 71% for much of the year except during summer (JJA) when it is in  
279 the 63%-66% range. The VCM displays a smaller difference between its error types with a TPR of 82%  
280 and a TNR of 92%, and a more limited seasonal dependence. These results show that VIIRS is able to  
281 strike a better balance between the Type I and Type II errors, while MAIAC's strength is its ability to  
282 greatly reduce false negatives in the AOT record, thereby reducing bias.

283 In terms of these Type I errors, since the MCM operates at both the block and pixel level, it is possible  
284 that diurnal convection produces sufficient cloud cover to cause the covariance between that block and  
285 the clear-sky reference image to decrease to the point that it is deemed cloudy. Likewise, cumulus cloud  
286 fields common over land during this season may be enough to trigger a cloudy designation for that pixel  
287 from MAIAC, while the very narrow field of view of the CALIOP sensor may pass between these small  
288 clouds leading to a conflicting collocation. Such instances of small clouds and sub-pixel clouds pose  
289 problems for all types of cloud masks produced by passive sensors.

290 Seasonal statistics (Fig. 2) showed that MAIAC has a significantly greater number of high quality  
291 retrievals than VIIRS in many U.S. regions, even those where the surface is not bright enough to keep  
292 the algorithm from performing the retrieval. This would imply that either MAIAC is opting to retrieve  
293 AOT in unfavorable conditions (presence of clouds/snow, etc.) or that VIIRS is failing to retrieve at a  
294 high quality over these areas. The results of the matchups with CALIPSO seem to suggest the later, as  
295 the MCM is being conservative in determining which pixels are cloud-free. Therefore, cloud screening  
296 is not thought to be a substantial driver behind the differences in retrieval numbers; however other limits  
297 placed on AOT retrievals within the algorithms may be playing a part in the spatial coverage.

298 Some recent preliminary analysis by the VIIRS Aerosol team into gaps in AOT over the CONUS has  
299 shown that the most probable cause for the reduced number of high quality IP retrievals is the limited

300 AOT range (0 to 2); and more precisely in this case, the lower bound of zero. Unlike VIIRS, which  
301 excludes the candidate pixel if the minimum residual corresponds to an AOT less than 0, MAIAC does  
302 not reject pixels whose surface reflectance falls below the expected value when computed with an AOD  
303 equal to 0. This happens on the occasion that the surface has changed significantly, or that the previous  
304 surface characterization is not correct. In the event this situation occurs, MAIAC reports an AOT of zero  
305 and then focuses on correcting the surface characterization with the next observation.

306 Large areas of missing AOT in VIIRS granules can be found in regions where the atmosphere is free of  
307 clouds or visible aerosols, meaning that the AOT is too small (negative) to be given a quality level high  
308 enough to be reported by the algorithm. This phenomenon is most prevalent in winter and spring when  
309 the AOT loading is small, and tends to be enhanced when the surface is sparsely vegetated and being  
310 viewed from the backscattering direction. In the recent VIIRS aerosol validation analysis performed by  
311 *Huang et al.*, (2016) it was shown that VIIRS is often negatively biased during the period from late fall  
312 to early spring. Additionally, *Liu et al.*, (2013) showed that VIIRS AOT tends to underestimate AOT  
313 when the surface is soil dominated. These two conclusions from previous validation studies support the  
314 notion that VIIRS has a tendency to retrieve more negative AOT when certain seasonal, geometric, and  
315 surface conditions are present, which can lead to relatively large areas with limited to no retrievals.

### 316 **3.2.3 Collocated retrievals of AOT**

317 As noted in the previous section, VIIRS and MAIAC tend to characterize the spatial patterns of seasonal  
318 AOT in similar ways. It also appears that MAIAC is generally a bit lower when compared to VIIRS,  
319 especially in the warm season. Observations collocated in time and space are needed to make sure that  
320 these two AOT products are being compared to one another under the same conditions. Therefore, the  
321 gridded data are filtered so that only days when both algorithms have enough retrievals to populate the  
322 grid cell are used in the analysis. Figure 3 presents the results of this collocation for the spring and

323 summer seasons when the differences between the two are greatest. While there is better agreement  
324 between MAIAC and VIIRS across much of the domain, the same trend of elevated AOT from VIIRS  
325 over the larger urban areas persists. Summer is the season with the highest disparity between the two  
326 algorithms, when a widespread difference between VIIRS and MAIAC is seen in the eastern half of the  
327 domain. In Figure 3d, this difference is shown to be predominately  $\pm 0.1$  or less; however there are small  
328 isolated pockets of larger bias up to 0.5. In other seasons, there is little systematic disagreement between  
329 the two with the exception of some high AOT from VIIRS over Montana and the Dakotas during the  
330 spring season. This discrepancy between the two could be a result of cloud contamination, or differences  
331 in surface characterization.

332 Those areas where VIIRS is significantly higher than MAIAC are likely caused by the underlying  
333 surface since many of these anomalies are predominately located over heavily urbanized areas and  
334 mountainous terrain. There are also smaller differences which are not as persistent but cover larger  
335 areas. An example of this can be seen in the summer season where VIIRS AOT in the eastern half of the  
336 U.S. is ubiquitously higher than MAIAC. Aerosol type and concentration can be widely different based  
337 on region, and problems characterizing these differences may be caused by certain underlying aspects of  
338 the aerosol algorithms.

339 One such component of the algorithms that could be responsible for the regional contrast is the different  
340 aerosol models used to retrieve AOT. MAIAC uses a dynamic model where physical parameters can  
341 change based on the magnitude of AOT. Volumetric concentrations of the fine and coarse particles can  
342 also be varied, thereby allowing for a wider range of size parameter to be simulated. In addition,  
343 MAIAC uses a background aerosol model that is tuned regionally based on AERONET optical thickness  
344 measurements. As a global product, VIIRS on the other hand uses five predefined aerosol models which  
345 have bimodal size distributions and static volumetric concentration parameters for each of the models  
346 and both particle sizes. Although not related to the aerosol models themselves, VIIRS also uses a

globally constant surface reflectance ratio to compare against the retrieved reflectance. This lack of accounting for such variations in surface type was discussed by *Liu et al.*, (2013) as a potential source of regional bias in the AOT retrievals. In that analysis it was also found that VIIRS is biased high in the Eastern U.S. when compared to both AERONET and MODIS. Together, these differences in aerosol models and surface characterization are capable of producing the regional variations in AOT retrieved from MAIAC and VIIRS.

### 3.3 Validation of products

#### 3.3.1 Comparison with AERONET AOT

In general, AOT from MAIAC and VIIRS compare well to one another, however there are differences and it is difficult to get a sense of which exhibits the higher level of accuracy without an ‘unbiased’ dataset to compare against. Measurements from AERONET sun photometers have been used for this purpose for many of the satellite derived aerosol products since the network’s inception [*Chu et al.*, 2002; *Kahn et al.*, 2005, *Liu et al.*, 2013]. Most recently, in a manuscript by *Huang et al.*, (2016) it was found using AERONET level 2 data that the VIIRS IP product has a global bias of 0.04. To determine the bias of the AOT produced by the two algorithms in question over our domain, we construct a set of matchups with AERONET level 2 data using the original datasets at their nominal resolution. *Petrenko et al.*, (2012) outlined a system for subsetting data from spaceborne sensors based on the location of ground-based sensors such as AERONET. This same process of matching our datasets with AERONET is used here, where all good quality retrievals within 27.5 km of the AERONET site are selected. As part of the matchup criteria, at least 20% of the total number of possible pixels within this circle are needed along with a minimum of 4 AERONET measurements over the time period of one hour centered on the satellite overpass time are required. All pixels found to meet these requirements are averaged together, as are all ground measurement that fall in the time window.



370 Figure 4 shows the scatter plots constructed using the AERONET matchups with VIIRS and MAIAC.  
 371 For all data matchups, VIIRS has a noticeable high bias which is pervasive at  $AOT < 0.04$ , and a  
 372 moderate correlation of 0.64 with AERONET. However, a VIIRS positive bias of 0.043 compares well  
 373 with the results of the global matchups presented in Huang et al., (2016). MAIAC on the other hand is  
 374 highly correlated (0.82) with AERONET and exhibits only a slight negative bias when compared with  
 375 AERONET. The greater number of MAIAC matchups is further evidence of its coverage and ability to  
 376 retrieve over the brighter surfaces over which many AERONET stations in the western U.S. are located.  
 377 In Figure 5, we highlight the dependence of the AOT bias on the magnitude of AOT by plotting the  
 378 differences between VIIRS and AERONET at 25 AOT bins of increasing size. The typical error  
 379 (median of all matchup errors) is often less than  $\pm 0.05$  with the exception of the strong negative bias for  
 380 both products during times of high aerosol loading, with MAIAC having slightly greater bias as AOT  
 381 increases. The spread of VIIRS errors however is much greater than those for MAIAC as evidenced by  
 382 the larger quartile ranges in most bins and the much higher maximum errors seen at low AOT.

383 Aerosol type is also an important consideration when evaluating the AOT retrievals since the chosen  
 384 aerosol model determines the spectral dependence of AOT. This spectral AOT can act as a proxy for  
 385 particle size, and the Angstrom Exponent (AE) is often used to qualitatively describe this spectral  
 386 dependence [Angstrom, 1929]. AE for coarse mode particles such as dust tend to be  $< 1$ , while finer  
 387 particles produced from urban pollution or biomass burning are associated with AE values  $> 2$  [Reid et  
 388 al., 1999; Schuster et al., 2006]. AERONET provides AE for multiple wavelength pairs and can be used  
 389 to determine if the retrieval errors from MAIAC or VIIRS are dependent on particle size. Figure 6  
 390 provides a look at how each algorithm performs across the range of particle sizes. The color coding of  
 391 the individual matchups is based on the AOT retrieved by AERONET. There is evidence of the larger  
 392 positive biases present and previously discussed in the VIIRS data which is limited to low-to-moderate  
 393 loading of finer particles. MAIAC meanwhile has very limited bias and dependence on particle size as

394 shown by the regression line. MAIAC however does have some issues retrieving accurately during high  
395 aerosol loading of coarse or mixed particle sizes (AE between 0.5 and 1.75). Figure 6 also reaffirms the  
396 results portrayed in Figure 4, however it shows that the larger biases tend to occur when the aerosol  
397 particle size is large, or when the concentration of coarse and fine particles is mixed. Both algorithms  
398 appear to perform quite well during cases of smoke or urban pollution.

399 While not analyzed directly here for reasons stated in Section 1, the demonstrated performance of the  
400 MODIS aerosol product is useful for providing extra context. A study from 2013 by Levy et al. details  
401 the performance of the MODIS Collection 6 algorithm and specifically section 4.4 outlines the MODIS  
402 Dark Target (DT) algorithm. Results for MAIAC shown here in Figures 4 & 5 compare well with the  
403 MODIS algorithm (Figure 11 in Levy et al.) over land with similar levels of accuracy and precision. It is  
404 important to note however that the Levy et al. study used global DT data, whereas MAIAC retrieves  
405 over both dark and bright surfaces and is constrained to the CONUS region in our analysis.

### 406 **3.3.2 Dependence of AOT on Viewing Geometry and Surface Reflectance**

407 In an attempt to ascertain which conditions might cause biases in the AOT retrievals, we look at how  
408 they are impacted by changing viewing geometry and surface brightness. Only data points where both  
409 VIIRS and MAIAC are matched with AERONET observations are used for this purpose, resulting in a  
410 dataset of 1034 matchups. Viewing geometry dependence is determined using the following 3  
411 parameters: viewing zenith angle; relative azimuth angle; and scattering angle. The AOT biases are  
412 separated into bins using 5 degree increments and plotted as a function of increasing angle. The results  
413 are shown in panels a, b, and c of Figure 7.

414 In terms of viewing angle, both algorithms produce matches that are well distributed across the range of  
415 angles with VIIRS having greater range as a result of the increased swath width over MODIS. MAIAC  
416 has very little viewing angle dependence, and has a minimal amount of a negative bias. VIIRS has some

417 viewing angle dependence with positive biases at-nadir that approach zero for larger VZA. The number  
418 of matchups are not as uniform for RAA, as both MAIAC and VIIRS have a bimodal distribution of  
419 angles with limited number of matches near  $90^\circ$ . MAIAC has some small dependence on RAA but  
420 biases are generally low except for the  $80\text{--}110^\circ$  range and near the extremes of  $0^\circ$  and  $180^\circ$  where  
421 matchups are very scarce. VIIRS AOT starts out with positive bias where strong back-scattering is  
422 occurring ( $\text{RAA} < 50^\circ$ ) with little dependency, however bias increases dramatically as the relative  
423 azimuth angle approaches  $180^\circ$ . It is worth noting that a limited amount of VIIRS matchups are  
424 available at  $\text{RAA} > 140^\circ$ , which is a range with both high bias and variability. Both algorithms have  
425 some bias dependence on scattering angle. MAIAC biases are within 0.02 of the zero line for smaller  
426 scattering angles, but the negative bias continues to get larger once SCA surpasses  $140^\circ$ . VIIRS also  
427 has a small negative bias which then becomes positive as scattering angle increases.

428 Figure 7d shows the dependence of the two algorithms in terms of the MAIAC surface reflectance which  
429 is binned at intervals of 0.005. Minimal errors are observed for both datasets over dark surfaces up to a  
430 reflectance of 0.06, after which the algorithms start slowly trending in different directions. The error  
431 becomes larger for VIIRS once the surface reflectance reaches 0.12, while MAIAC dependence on  
432 surface reflectance reverses after this point. The brighter surfaces also appear to cause increased  
433 fluctuation in bias for both of the algorithms

434 As noted previously, there is some dependence on sun-sensor geometry for both of the algorithms  
435 analyzed here. Notably, there is a large difference in the level of dependence between retrievals in the  
436 back-scattering direction ( $\text{RAA} < 90^\circ$ ) and the forward-scattering direction ( $\text{RAA} > 90^\circ$ ) for VIIRS. The  
437 two algorithms also drift away from the zero line in opposite directions for scattering angles greater than  
438  $100^\circ$ . Due to the anisotropy of surface reflectance for many land targets, this change in viewing direction  
439 can lead to changes in the perceptible brightness of the surface, a phenomenon known as directional  
440 scattering. This effect causes an apparent brightening of the surface when viewed from in the back-

441 scattering direction, and some dimming in the forward-scattering direction [Roujean, 1992]. MAIAC,  
442 through its use of the BRDF when retrieving AOT, attempts to account for and mitigate these effects.  
443 Based on the results in Figure 7d, it appears as though it is able to remove much of this dependence;  
444 VIIRS meanwhile, because of the assumption of a Lambertian surface, produces AOT with higher  
445 biases.

446 To see how each algorithm handles these changes the matchups for surface reflectance have been further  
447 stratified based on the scattering direction (using RAA of  $90^\circ$  as a separator). The resulting biases and  
448 histograms for both directions are given in Figure 8. VIIRS dependencies are similar regardless of the  
449 scattering direction, although errors are markedly higher in the forward-scattering direction for brighter  
450 surfaces. On the other hand, the dependency for MAIAC does look quite different depending on the  
451 scattering direction. MAIAC errors are near zero over dark surfaces in the back-scattering direction, yet  
452 quickly become negative as the surface gets brighter. In the forward-scattering direction, a rather  
453 consistent negative bias around -0.05 is found until surface reflectance surpasses 0.12, when it becomes  
454 more varied. Comparing these two panels to Figure 6d, we see that the back-scattering retrievals tend to  
455 dominate the overall signal due to nearly two-thirds of the retrieval matchups falling within this relative  
456 azimuth range; with the only exception being the bright surfaces where MAIAC has few valid retrievals.  
457 The histograms also show that MAIAC has some offset in the surface reflectance of its retrievals in both  
458 directions when compared to VIIRS. This is likely a result of including the BRDF in its retrieval strategy  
459 which accounts for the effects of sun and satellite geometry thereby reducing the brightness in the  
460 backward direction and increasing it in the forward direction.

### 461 **3.3.3 Sources of Bias**

462 Matchups of MAIAC and VIIRS with AERONET data in the U.S. and surrounding areas have shown  
463 that biases are present that are angular dependent. MAIAC dependencies are less pronounced than  
464 VIIRS, but a negative association with geometric surface attributes does exist. *Lyapustin et al., (2011)*

465 showed that SRC does vary slightly with viewing geometry, and that the use of an average SRC value  
466 will cause the algorithm to overestimate surface reflectance in the forward direction and vise-versa for  
467 back-scattering geometries. This reduced brightening in the backward direction and increase in the  
468 surface reflectance in the forward scattering direction relative to VIIRS is evident in the histogram  
469 offsets seen in Figure 8.

470 The consequence of this would be an underestimated AOT in the forward-scattering direction, and  
471 overestimation in the back-scattering direction, however we only find a consistent negative bias in the  
472 forward direction. In the back-scattering direction, the surface tends to be brighter due to reduced  
473 shadowing and lower aerosol backscattering compared to the forward-scattering direction. This can  
474 cause the sensitivity of the TOA reflectance to AOT to decrease, leading to higher uncertainty of AOT  
475 in the back-scattering direction. This combined with the limited amount of MAIAC matchups with a  
476 high surface reflectance in the back-scattering direction are likely leading to the larger, variable errors  
477 over bright surfaces.

478 Previous global validation studies have focused on VIIRS Aerosol products [*Liu et al.*, 2013 (EDR  
479 only); *Huang et al.*, 2016 (EDR and IP)] and have shown that a slight positive bias is observed in AOT  
480 over land. As is shown in this analysis, *Liu et al.*, (2013) also found a similar dependence in the EDR  
481 data in relationship to viewing zenith angle over land as is shown in this paper, although errors were  
482 found to be larger in this case. This is not surprising as more noise is expected in the pixel-level IP AOT  
483 data, which does not have the benefit of aggregation and further screening. Even with that in mind, the  
484 level of bias seen in this study for VIIRS products is concerning since data at this product level is useful  
485 to the air quality community who require highly accurate data for their applications. Therefore, a brief  
486 attempt was made to uncover additional sources of bias to those already established by previous studies.

Recall from section 3.2 that urban ‘hotspots’ of AOT were consistently present over medium to large cities across the U.S. in all seasons (more so in warm seasons). A fair amount of AERONET sites that are not surrounded by bright or soil-dominated surfaces in the U.S. are located in or near these urban areas, meaning that some of the bias may be attributed to these sites. In fact, of those matchups which exhibit excessive positive bias ( $> 0.1$ ), 65% of them are associated with a handful of sites located in Los Angeles or Houston, two large and highly urbanized cities. Over 85% of the highly biased matchups (20% of all matchups) originate from AERONET sites located in a major metro area. When looking at viewing geometry values where large biases are seen, we notice a considerable number of those AERONET sites also being in select urban areas, while sites with lower biases tend to be more random. This suggests that a sizable portion of the large biases and dependencies on viewing geometry in this domain may be due to a lack of accuracy over urban areas and that viewing geometry is an intensifier of those biases.

#### **4. AOT case studies**

Up to this point, the geographic inspection of the AOT products from MAIAC and VIIRS have been contained to seasonally gridded AOT. In an attempt to observe and verify some of the findings from the bias analysis, a look at individual cases at the products’ native resolution are presented below. This allows for qualitative comparison of the two products independent of the AERONET matchups which, with respect to VIIRS, were found to be heavily influenced by an urban bias. Two cases; one with a large area of smoke present over the northwestern U.S. and a more typical late-summer AOT case in the eastern half of the country were chosen. Careful attention was paid to make certain that the Aqua and Suomi-NPP overpass times for the selected date were close together ( $< 20$  min) so valid spatial comparisons could be made.

##### **4.1 High AOT case**

510 In 2013, a few large historical wildfires took place in North America with one such fire being the Rim  
511 fire, which started on August 17th near Yosemite National Park and burned for over two months. Figure  
512 9 shows a VIIRS true-color image over the Western U.S. from August 25<sup>th</sup> along with AOT from VIIRS  
513 and MAIAC. The two products agree well over regions where both have retrieved AOT, however  
514 differences do exist. VIIRS IP AOT is higher over the thickest parts of the smoke plume and is noisier,  
515 however this is expected since it is a pixel-level product while the MAIAC AOT has the advantage of  
516 using gridded MODIS reflectance, and much of the information used to perform the retrieval is supplied  
517 from processing at the block-level.

518 Just as the analysis in section 3 showed, VIIRS coverage over brighter surfaces is limited compared to  
519 MAIAC, as large sections of Montana, Idaho and Wyoming lack any high quality retrievals. However  
520 VIIRS does retrieve more of the smoke in northern Idaho. The missing MAIAC retrievals in the far  
521 upper right section of the image are a result of it being outside MAIAC's North American processing  
522 domain. There are also smaller rectangular holes in the MAIAC data near the center of the image which  
523 are a product of the block-level SRC retrieval that takes place within the aerosol retrieval. In some  
524 cases, SRC may not be retrieved or updated due to cloudiness. This causes AOT to not be retrieved over  
525 the brighter surfaces within that block (25 km x 25km).

## 526 **4.2 Moderate AOT case**

527 Given that strong AOT bias dependencies exist in both the viewing geometry and AOT itself, a second  
528 case representing a more moderate aerosol loading scenario was investigated. Figure 10 includes the  
529 true-color image and AOT maps from VIIRS and MAIAC on Sept. 5<sup>th</sup>, 2013 over the Mid-western and  
530 Mid-Atlantic states. In contrast to the previous example, the spatial coverage of VIIRS is much closer to  
531 MAIAC in this case due to a majority of the surface being dark. The exception here is over inland water  
532 bodies such as the Great Lakes where VIIRS currently does not retrieve AOT. Once again, the two

533 products characterize the spatial variation in AOT in similar ways. Over much of the Ohio River valley,  
534 where an area of haze exists, the two algorithms produce results that are very alike, although VIIRS  
535 AOT is slightly higher in the vicinity of clouds in northern Illinois. VIIRS is also significantly higher  
536 over the Chicago and St. Louis urban areas which are circled in black, lending credibility to the theory  
537 that VIIRS is often biased high over cities. No AERONET sites are located in Chicago, but one is  
538 located in downtown St. Louis, where data shows that VIIRS is biased high by 0.05 while MAIAC has a  
539 bias of -0.11. There are also areas where VIIRS is retrieving slightly higher AOT in a more uniform  
540 manner. The clearest example of this is in the Mid-Atlantic where VIIRS is retrieving AOTs which are  
541 around 0.05 higher than MAIAC. A similar pattern is also visible over a region stretching from Lake  
542 Michigan into Ohio and Pennsylvania.

543

## 544 **5. Conclusions**

545 This study was undertaken to assess the utility of the MAIAC algorithm for retrieving aerosol  
546 information from a passive satellite sensor through a comparison with the aerosol products from VIIRS  
547 and ground-based sun photometers. With these data sets as benchmarks, we were able to evaluate the  
548 spatial coverage and accuracy of the MAIAC AOT product. Using data gridded to 0.25 degrees, we  
549 found that MAIAC is capable of providing retrievals over a varied set of surface types, including the  
550 bright and soil dominated surfaces which restrict the coverage of the common dark-target only  
551 algorithms (VIIRS, MODIS). The number of valid high-quality retrievals MAIAC produces is also  
552 greater, leading us to evaluate the cloud mask performance of both algorithms through matchups with  
553 CALIOP. Those matchups showed that both MAIAC and VIIRS had similar accuracy, however we  
554 found MAIAC to be more conservative in its assignment of clear-sky pixels. When compared directly  
555 with VIIRS, MAIAC produces AOT values that on average are 0.017 lower than VIIRS during 2013.



556 There is large seasonality however, with minor differences for winter and fall, and larger separation seen  
557 in the summer season.

558 In order to conduct a more robust accuracy assessment including the dependence of the algorithms on  
559 viewing geometry and surface reflectance, both datasets were also evaluated against AERONET Level 2  
560 AOT. MAIAC showed little dependence on viewing zenith, however there was some negative  
561 association with the scattering angle and the brightness of the surface. VIIRS showed negative  
562 association with viewing angle, but was positive with scattering angle and surface reflectance. Biases as  
563 a function of surface reflectance were further stratified based on scattering direction because of the  
564 differences in errors seen with both products. Trends in VIIRS bias as a function of surface reflectance  
565 were not greatly affected by scattering direction, although overall errors were larger in the forward-  
566 scattering direction. Analysis of MAIAC showed that it only has strong dependence on surface  
567 reflectance when the surface is viewed in the back-scattering direction.

568 The results of this bias analysis coincided well with the initial investigations of the MAIAC algorithm.  
569 The results after studying the VIIRS biases with respect to scattering direction however were not  
570 consistent with previous validation studies; therefore a closer look was taken at those highly biased  
571 matchups. It was found that urban backgrounds may be causing, or at least intensifying the positive bias  
572 seen in VIIRS AOT. Overall, the MAIAC algorithm has shown the ability to perform well over the  
573 North American region with a high level of accuracy given its spatial resolution. Global analysis over a  
574 longer time period will be needed to make certain that the product(s) are robust and meet the levels of  
575 accuracy needed for aerosol monitoring.

576

577 **Acknowledgements.** The work outlined in this manuscript was supported by the NOAA JPSS program  
578 office and NASA as part of a GEO-CAPE aerosol science study. In addition, the authors would like to

579 recognize the following for their support throughout the study: I. Laszlo, H. Liu along with the entire  
580 NOAA/NESDIS/STAR Aerosol team. We also thank the AERONET principle investigators and their  
581 staff for the establishment and maintenance of the North American sites used as part of this  
582 investigation. VIIRS aerosol products are available at NOAA's Comprehensive Large Array-Data  
583 Stewardship System (CLASS, <http://www.class.ngdc.noaa.gov/>) MAIAC Aerosol products can be  
584 accessed through at the NASA ftp site:  
585 <ftp://maiac@dataportal.nccs.nasa.gov/DataRelease/NorthAmerica>. The AERONET data used in this  
586 study is publically available from the AERONET team at the NASA GFSC site  
587 (<http://aeronet.gsfc.nasa.gov/>). Data from CALIPSO were obtained from the NASA Langley Research  
588 Center Atmospheric Science Data Center (ASDC). The contents of this manuscript are solely the  
589 opinions of the authors and do not constitute a statement of policy, decision, or position on behalf of  
590 NOAA or the U.S. Government.

591

592

593

## 594 **References**

- 595 Al-Saadi, J., Szykman, J., Pierce, R. B., Kittaka, C., Neil, D., Chu, D. A., & Fishman, J. (2005).  
596 Improving national air quality forecasts with satellite aerosol observations. *Bulletin of the American*  
597 *Meteorological Society*, 86(9), 1249-1261.
- 598 Ångström, A. (1929). On the atmospheric transmission of sun radiation and on dust in the  
599 air. *Geografiska Annaler*, 11, 156-166.

Chu, D. A., Kaufman, Y. J., Ichoku, C., Remer, L. A., Tanré, D., & Holben, B. N. (2002). Validation  
 of MODIS aerosol optical depth retrieval over land. *Geophysical research letters*, 29(12), MOD2-1.

Fishman, J., L. T. Iraci, J. Al-Saadi, Kelly V. Chance, F. Chavez, M. Chin, P. Coble, 2012. “The  
 United States’ Next Generation of Atmospheric Composition and Coastal Ecosystem Measurements:  
 NASA’s Geostationary Coastal and Air Pollution Events (GEO-CAPE) Mission.” *Bulletin of the  
 American Meteorological Society* 93 (10) (October): 1547–1566. doi:10.1175/bams-d-11-00201.1.

Frey, R. A., Ackerman, S. A., Liu, Y., Strabala, K. I., Zhang, H., Key, J. R., & Wang, X. (2008).  
 Cloud detection with MODIS. Part I: Improvements in the MODIS cloud mask for collection  
 5. *Journal of Atmospheric and Oceanic Technology*, 25(7), 1057-1072.

Gao, B. C., & Kaufman, Y. J. (2003). Water vapor retrievals using Moderate Resolution Imaging  
 Spectroradiometer (MODIS) near-infrared channels. *Journal of Geophysical Research:  
 Atmospheres*, 108(D13).

Heidinger, A. K., Evan, A. T., Foster, M. J., & Walther, A. (2012). A naive Bayesian cloud-  
 detection scheme derived from CALIPSO and applied within PATMOS-x. *Journal of Applied  
 Meteorology and Climatology*, 51(6), 1129-1144.

Holben, B., Vermote, E., Kaufman, Y. J., Tanré, D., & Kalb, V. (1992). Aerosol retrieval over land  
 from AVHRR data-application for atmospheric correction. *Geoscience and Remote Sensing, IEEE  
 Transactions on*, 30(2), 212-222.

Holben, B. N., Eck, T. F., Slutsker, I., Tanre, D., Buis, J. P., Setzer, A., & Smirnov, A. (1998).  
 AERONET—A federated instrument network and data archive for aerosol characterization. *Remote  
 sensing of environment*, 66(1), 1-16.

Huang, J., S. Kondragunta, I. Laszlo, H. Liu, L. A. Remer, H. Zhang, S. Superczynski, P. Ciren, B. N. Holben, and M. Petrenko (2016), Validation and expected error estimation of Suomi-NPP VIIRS aerosol optical thickness and Ångström exponent with AERONET, *Journal of Geophysical Research: Atmospheres*, 121, doi:10.1002/2016JD024834.

Holben B., & Petrenko M. (2016). Validation and Expected Error Estimation of S-NPP VIIRS Aerosol Optical Thickness and Angstrom Exponent with AERONET. *Journal of Geophysical Research: Atmospheres*, 121, doi:10.1002/2016JD024834.

Jackson, J. M., Liu, H., Laszlo, I., Kondragunta, S., Remer, L. A., Huang, J., & Huang, H. C. (2013). Suomi-NPP VIIRS aerosol algorithms and data products. *Journal of Geophysical Research: Atmospheres*, 118(22), 12-673.

Kahn, R. A., Gaitley, B. J., Martonchik, J. V., Diner, D. J., Crean, K. A., & Holben, B. (2005). Multi-Angle Imaging Spectroradiometer (MISR) global aerosol optical depth validation based on 2 years of coincident Aerosol Robotic Network (AERONET) observations. *Journal of Geophysical Research: Atmospheres (1984–2012)*, 110(D10).

Kaufman, Y. J., Tanré, D., Remer, L. A., Vermote, E. F., Chu, A., & Holben, B. N. (1997). Operational remote sensing of tropospheric aerosol over land from EOS moderate resolution imaging spectroradiometer. *Journal of Geophysical Research: Atmospheres (1984–2012)*, 102(D14), 17051-17067.

Kopp, T. J., Thomas, W., Heidinger, A. K., Botambekov, D., Frey, R. A., Hutchison, K. D., & Reed, B. (2014). The VIIRS Cloud Mask: Progress in the first year of S-NPP toward a common cloud detection scheme. *Journal of Geophysical Research: Atmospheres*, 119(5), 2441-2456.

Lahoz, W. A., Peuch, V. H., Orphal, J., Attié, J. L., Chance, K., Liu, X., ... & Amraoui, L. E. (2012). Monitoring air quality from space: The case for the geostationary platform. *Bulletin of the American Meteorological Society*, 93(2), 221-233.

Liu, H., Remer, L. A., Huang, J., Huang, H. C., Kondragunta, S., Laszlo, I., & Jackson, J. M. (2014). Preliminary evaluation of S-NPP VIIRS aerosol optical thickness. *Journal of Geophysical Research: Atmospheres*, 119(7), 3942-3962.

Lyapustin, A., Wang, Y., & Frey, R. (2008). An automatic cloud mask algorithm based on time series of MODIS measurements. *Journal of Geophysical Research: Atmospheres* (1984–2012), 113(D16).

Lyapustin, A., Martonchik, J., Wang, Y., Laszlo, I., & Korkin, S. (2011). Multi-Angle implementation of atmospheric correction (MAIAC): 1. Radiative transfer basis and look-up tables. *Journal of Geophysical Research: Atmospheres* (1984–2012), 116(D3).

Lyapustin, A., Wang, Y., Laszlo, I., Kahn, R., Korkin, S., Remer, L., & Reid, J. S. (2011). Multi-Angle implementation of atmospheric correction (MAIAC): 2. Aerosol algorithm. *Journal of Geophysical Research: Atmospheres* (1984–2012), 116(D3).

Lyapustin, A. I., Wang, Y., Laszlo, I., Hilker, T., Hall, F. G., Sellers, P. J., & Korkin, S. V. (2012). Multi-angle implementation of atmospheric correction for MODIS (MAIAC): 3. Atmospheric correction. *Remote Sensing of Environment*, 127, 385-393.

Lyapustin, A., M. J. Alexander, L. Ott, A. Molod, B. Holben, J. Susskind, and Y. Wang (2014), Observation of mountain lee waves with MODIS NIR column water vapor, *Geophysical Research Letters*, 41, doi:10.1002/2013GL058770

NOAA-NESDIS (2014), Joint Polar Satellite System (JPSS) Program Level 1 Requirements Document Supplement JPSS-REQ-1002. Version 2.10 [Retrieved from [http://www.jpss.noaa.gov/assets/pdfs/technical\\_documents/level\\_1\\_requirements\\_supplement.pdf](http://www.jpss.noaa.gov/assets/pdfs/technical_documents/level_1_requirements_supplement.pdf)]

NRC (2007): Earth Science and Applications from Space: National Imperatives for the Next Decade and Beyond. The National Academies Press, 400 pp.

Petrenko, M., Ichoku, C., & Leptoukh, G. (2012). Multi-sensor aerosol products sampling system (MAPSS). *Atmospheric Measurement Techniques*, 5(5), 913-926.

Ramanathan, V. C. P. J., Crutzen, P. J., Kiehl, J. T., & Rosenfeld, D. (2001). Aerosols, climate, and the hydrological cycle. *Science*, 294(5549), 2119-2124.

Reid, J. S., Eck, T. F., Christopher, S. A., Hobbs, P. V., & Holben, B. (1999). Use of the Ångström exponent to estimate the variability of optical and physical properties of aging smoke particles in Brazil. *Journal of Geophysical Research: Atmospheres*, 104(D22), 27473-27489.

Roujean, J. L. (1992). for the Correction of Remote Sensing Data. *Journal of geophysical research*, 97(D18), 20-455.

Schuster, G. L., Dubovik, O., & Holben, B. N. (2006). Angstrom exponent and bimodal aerosol size distributions. *Journal of Geophysical Research: Atmospheres*, 111(D7).

Smirnov A., B.N.Holben, T.F.Eck, O.Dubovik, and I.Slutsker, 2000: Cloud screening and quality control algorithms for the AERONET database, *Rem.Sens.Env.*, **73**, 337-349.

Stephens, G. L., Vane, D. G., Boain, R. J., Mace, G. G., Sassen, K., Wang, Z., ... & CloudSat Science Team, T. (2002). The CloudSat mission and the A-Train: A new dimension of space-based observations of clouds and precipitation. *Bulletin of the American Meteorological Society*, 83(12), 1771-1790.

Tanré, D., Kaufman, Y. J., Herman, M., & Mattoo, S. (1997). Remote sensing of aerosol properties over oceans using the MODIS/EOS spectral radiances. *Journal of Geophysical Research: Atmospheres* (1984–2012), 102(D14), 16971-16988.

Vaughan, M. A., Young, S. A., Winker, D. M., Powell, K. A., Omar, A. H., Liu, Z., & Hostetler, C. A. (2004, November). Fully automated analysis of space-based lidar data: An overview of the CALIPSO retrieval algorithms and data products. In *Remote Sensing* (pp. 16-30). International Society for Optics and Photonics.

Vermote, E. F., & Kotchenova, S. (2008). Atmospheric correction for the monitoring of land surfaces. *Journal of Geophysical Research: Atmospheres* (1984–2012), 113(D23).

Wang, J., & Christopher, S. A. (2003). Inter-comparison between satellite-derived aerosol optical thickness and PM<sub>2.5</sub> mass: implications for air quality studies. *Geophysical research letters*, 30(21).

Winker, D. M., Vaughan, M. A., Omar, A., Hu, Y., Powell, K. A., Liu, Z, Hunt, W. H., Young, S. A. (2009). Overview of the CALIPSO mission and CALIOP data processing algorithms. *Journal of Atmospheric and Oceanic Technology*, 26(11), 2310-2323.

## Tables

Table 1. Confusion matrix showing the designation of pixels from each cloud mask associated with the two algorithms compared with information on clouds from CALIPSO lidar taken as the “truth” datasets. The abbreviations in parenthesis note the location of the following test outcomes for both sets of data: True Positive (TP); False Positive (FP); False Negative (FN); and True Negative (TN).

		VIIRS		MAIAC	
		Cloudy	Clear	Cloudy	Clear
CALIPSO	Cloudy	65079 (TP)	14479 (FN)	1055111	40781
	Clear	4129 (FP)	47298 (TN)	235293	605130

Accuracy	86%	86%
----------	-----	-----

706

707

708

709

710

711

712

713

714

715

716

717

718

719

720

721

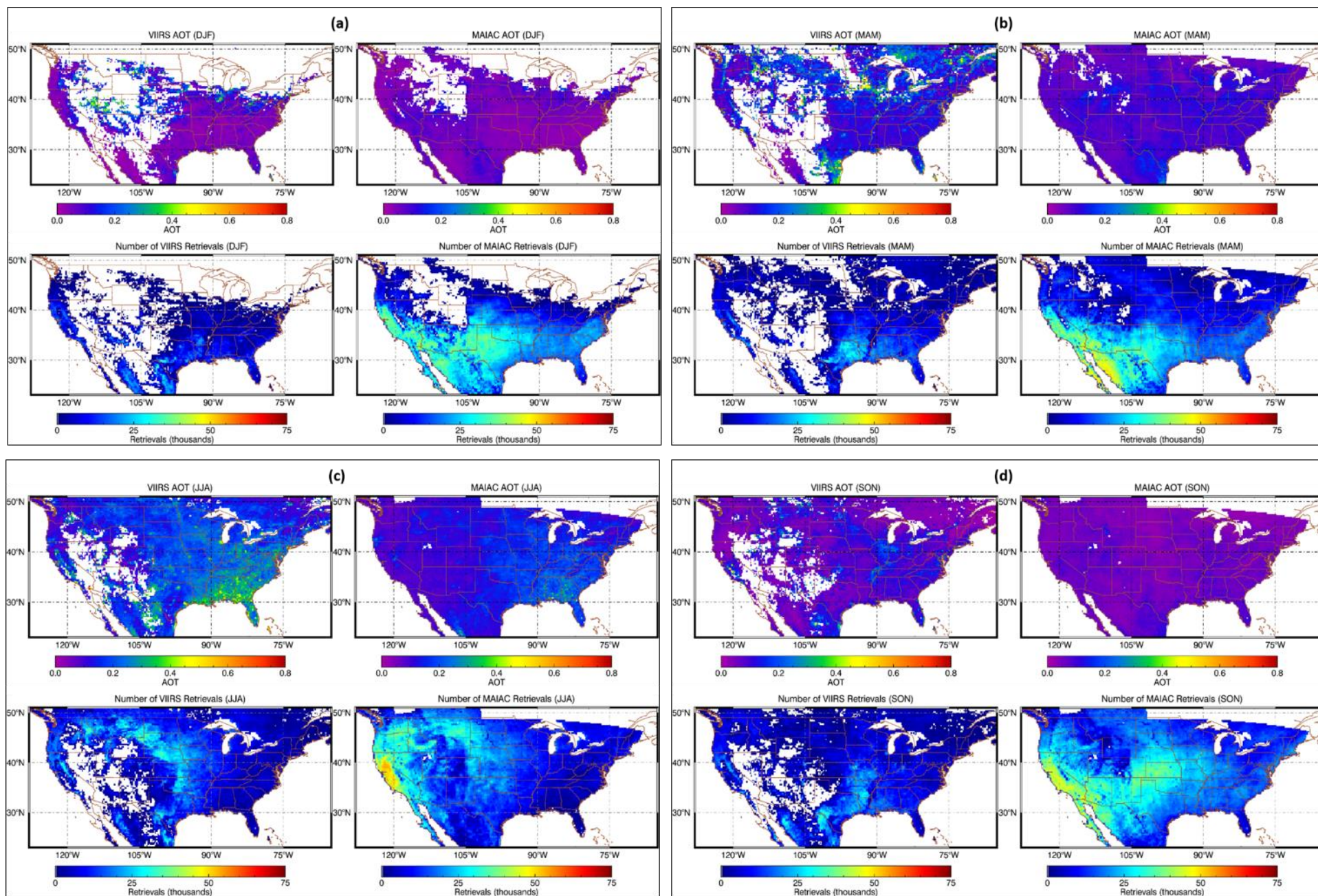
722

**Figures**





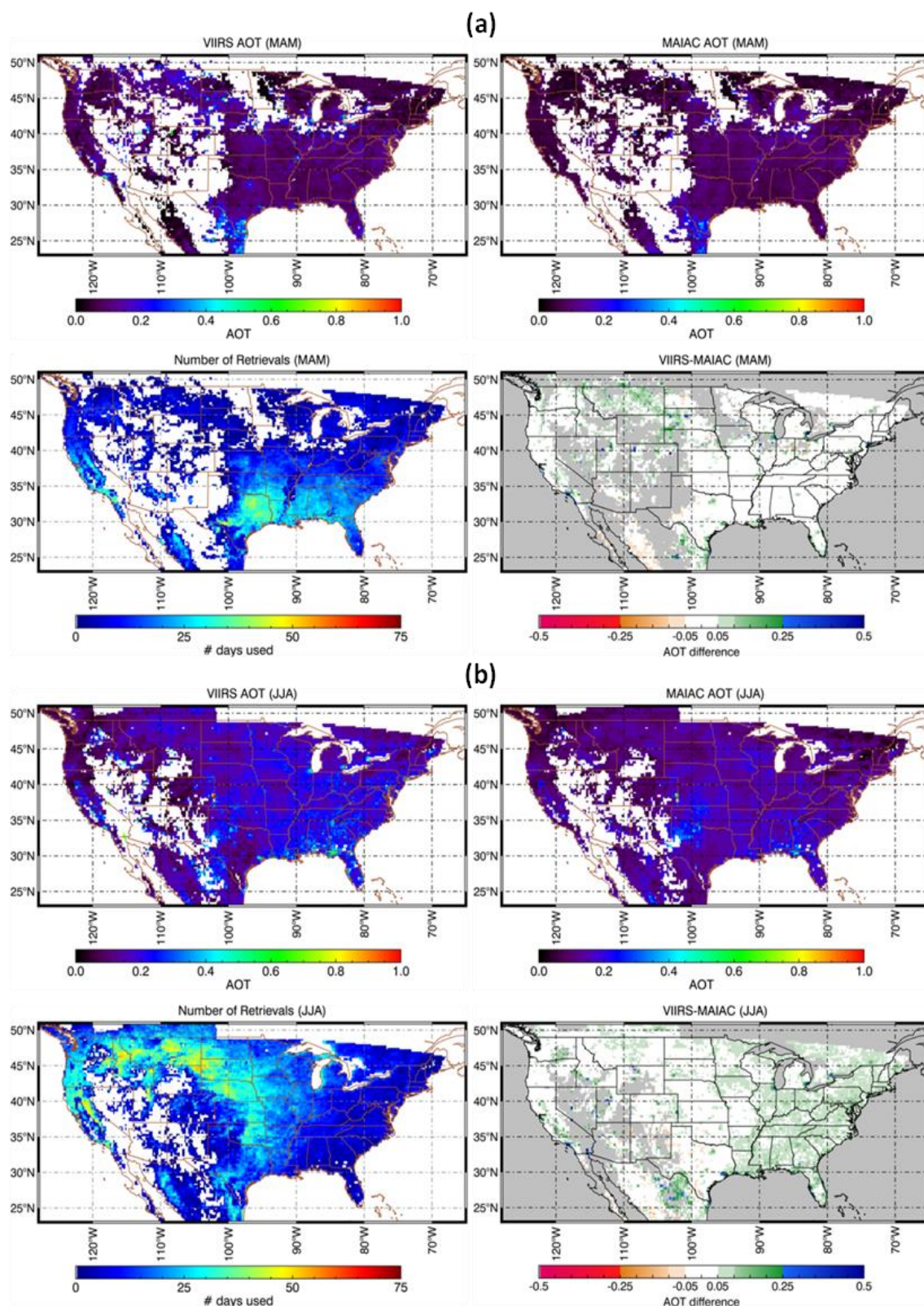
Figure 1. Map of the domain area used to grid the MAIAC and VIIRS AOT datasets. Domain was chosen based on the extent of MAIAC data currently available over North America. Coordinates of the upper left corner are (51° N, 129° W) and the lower right coordinates are (22° N, 65° W). Map data courtesy of Google Earth Pro (V 7.1.2.2041), Landsat.



730 Figure 2. Maps of gridded AOT at 550 nm (top) and retrieval count (bottom) from VIIRS and MAIAC for: (a) winter; (b) spring; (c) summer;  
731 and (d) fall. Large portions of missing data in MAIAC maps over southern Canada are caused by the geographic extent of available data in  
732 this region.

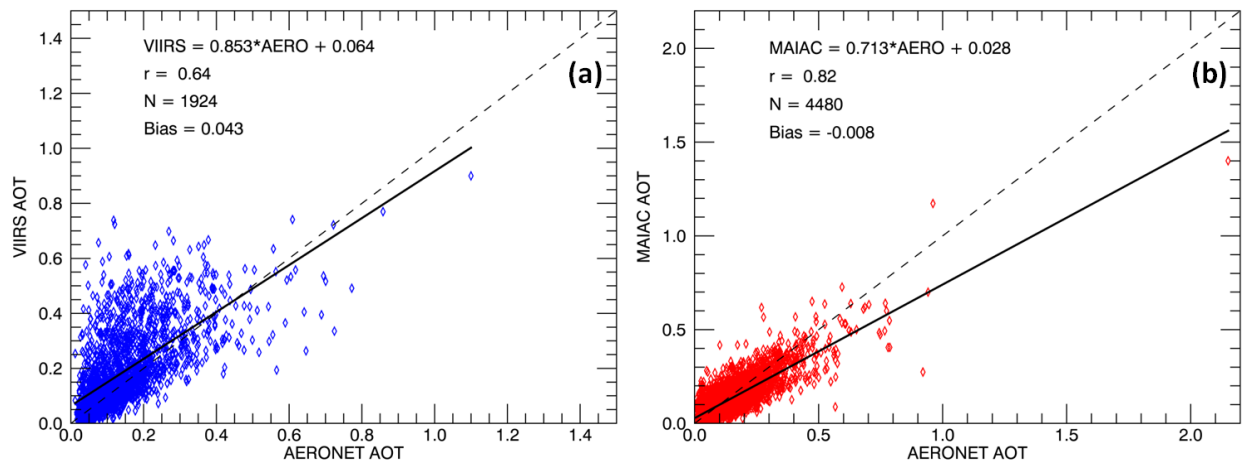






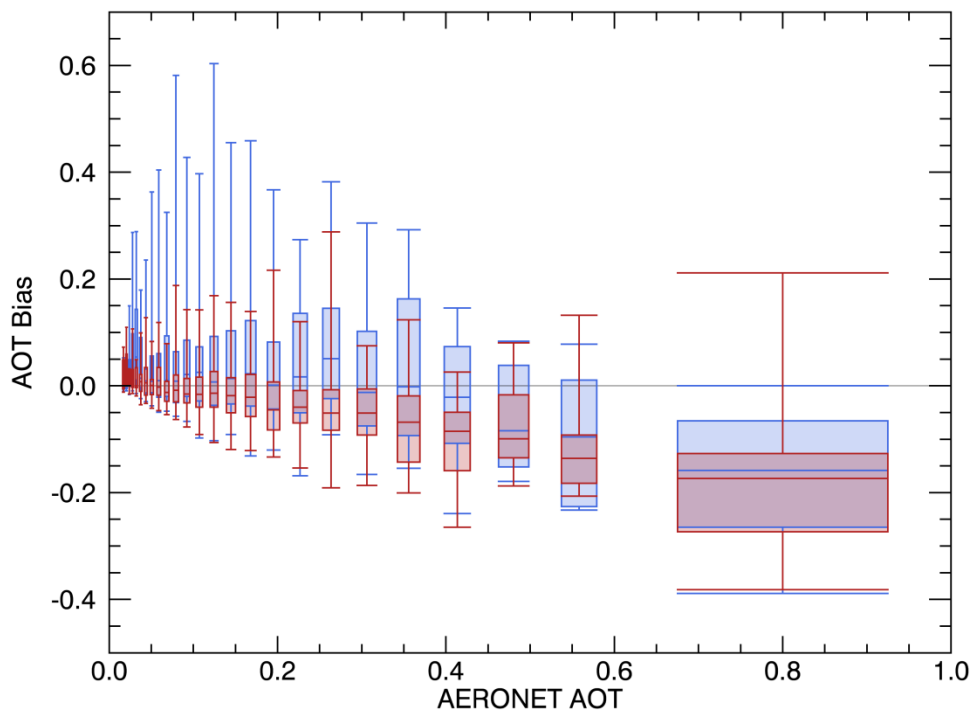
734

735 Figure 3. Set of four-panel plots showing matched VIIRS (upper left) and MAIAC (upper right) AOT  
 736 along with number of days with coincident observations (lower left), and AOT difference between the  
 737 products (lower right) for the spring (a) and summer (b) seasons.



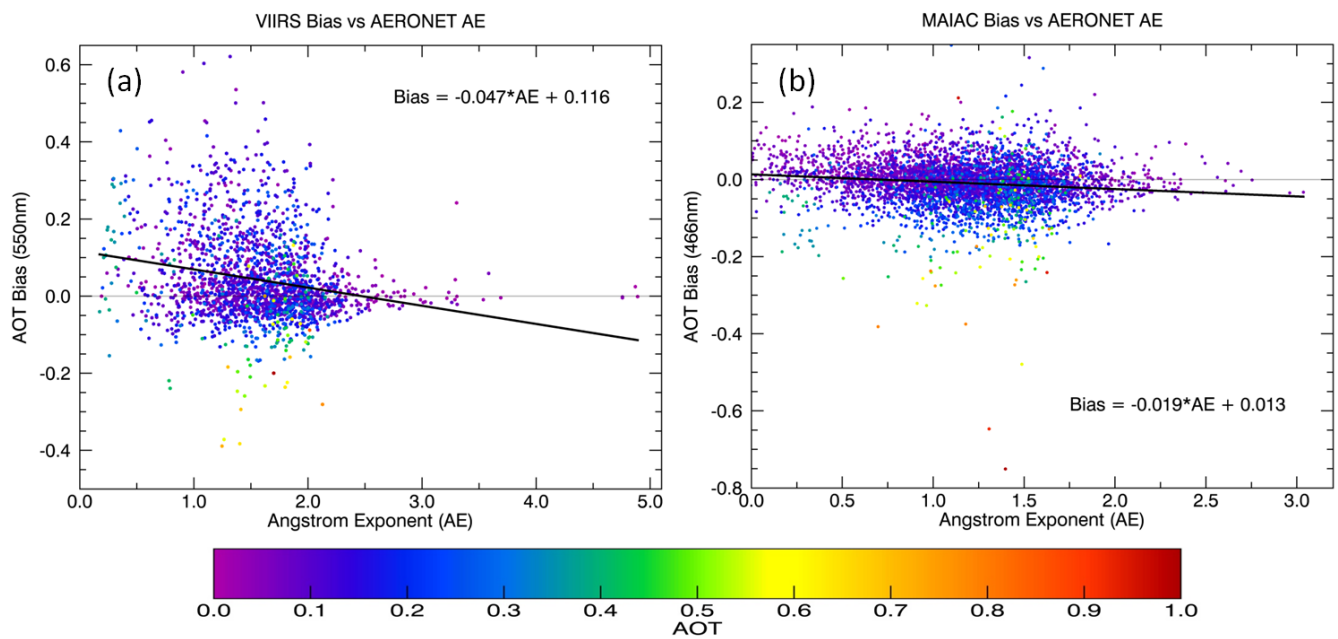
739

740 Figure 4. Scatter plots showing the relationship between AERONET AOT and VIIRS (a) and MAIAC  
741 (b). The dashed line represents the 1:1 line where the two datasets would be in complete agreement,  
742 while the solid lines represent the linear regression model (chi-squared test) provided at the top of each  
743 figure. Relevant relational statistics for correlation,  $r$ ; number of observation,  $N$ ; and bias are also given.  
744



745

746 Figure 5. Box and whisker plot showing the dependence of the VIIRS (blue) and MAIAC (red) bias on  
747 the AOT as measured by AERONET. Any missing data is due to the lack of matchups ( $< 5$ ) in that  
748 AOT bin.



749

750 Figure 6. AOT errors from the (a) VIIRS, and (b) MAIAC matchups as a function of AERONET  
 751 Angstrom Exponent, with regression line drawn in black. Data points are color-coded based on the  
 752 AERONET AOT retrieval associated with those matchups.

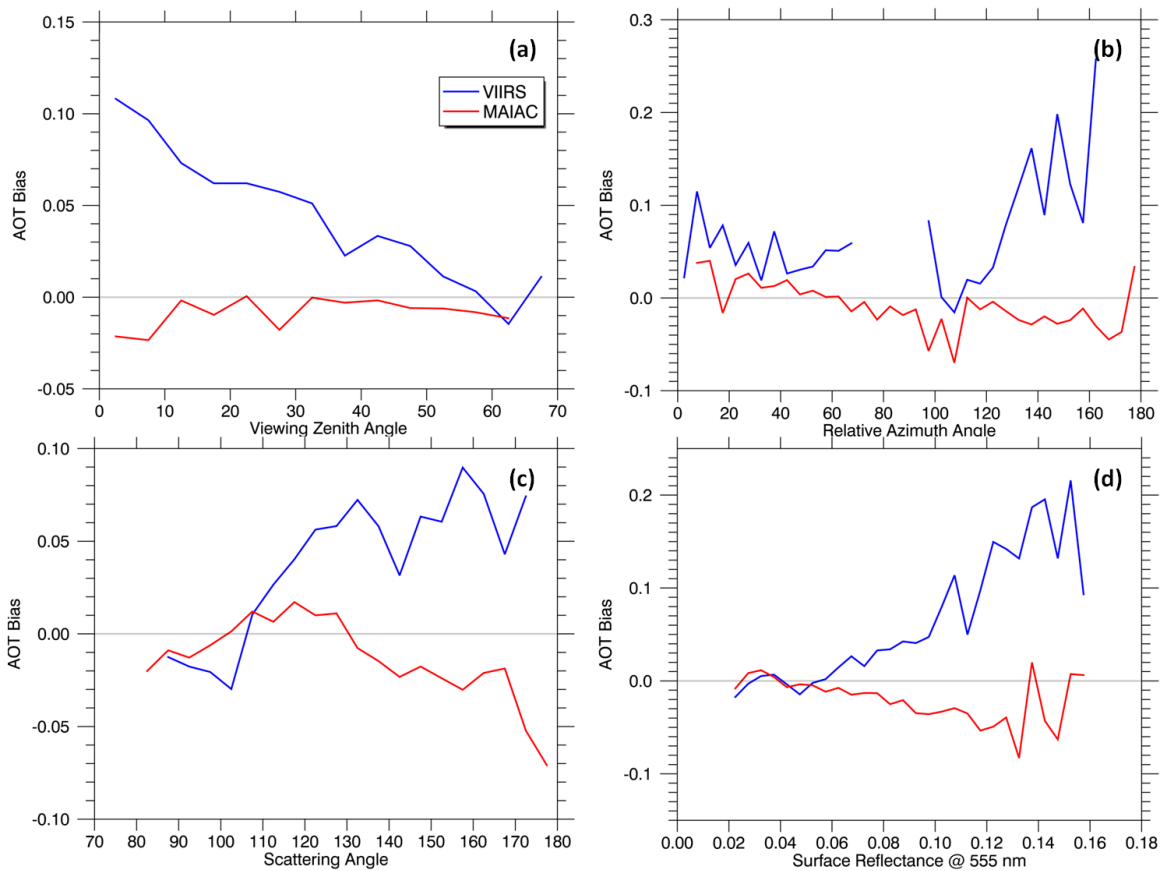


Figure 7. Dependence of AOT bias on: (a) viewing zenith angle; (b) relative azimuth angle; (c) scattering angle; and (d) surface reflectance at 555 nm according to MAIAC. VIIRS data is shown in blue and MAIAC in red, while the horizontal zero line (gray) is added for reference.

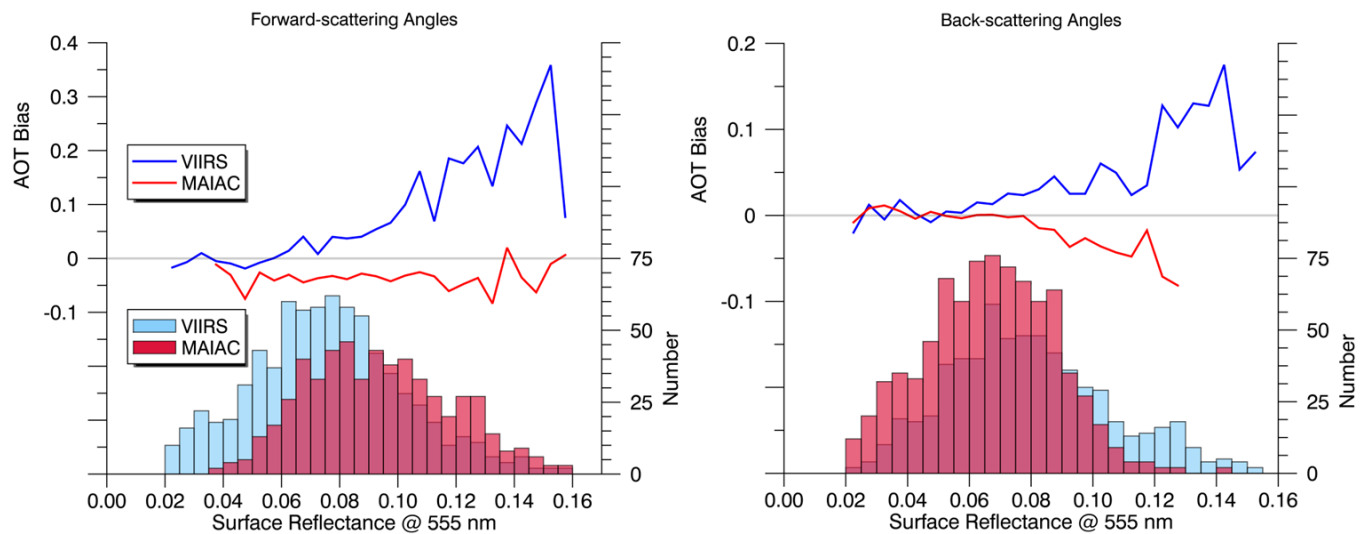
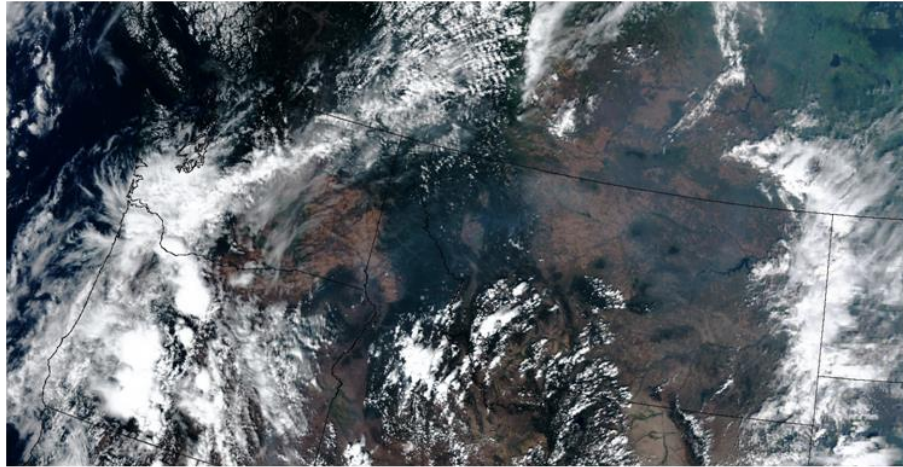


Figure 8. Dependencies on surface brightness split into observations taken from the forward-scattering (left) and back-scattering (right) direction. Bias is on the left-hand vertical axis and represented by the

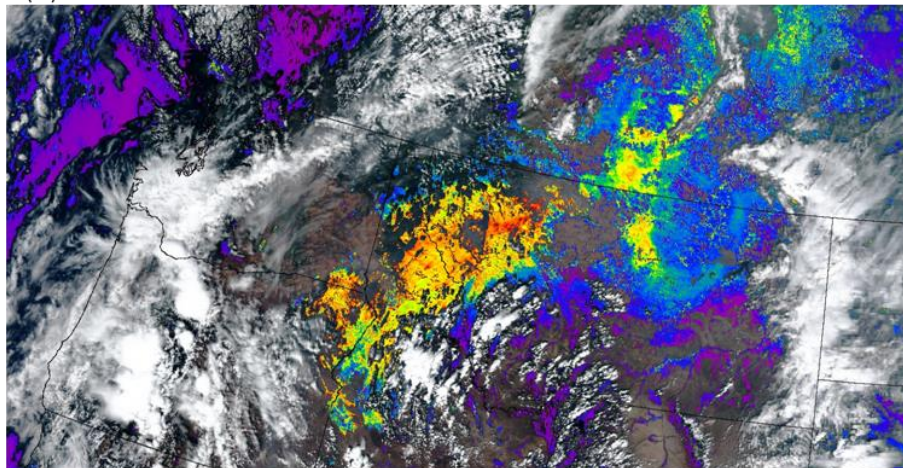


761 vertical lines, while the number of matchups in each reflectance bin are given by the vertical bars and  
762 occupy the right-hand vertical axis.

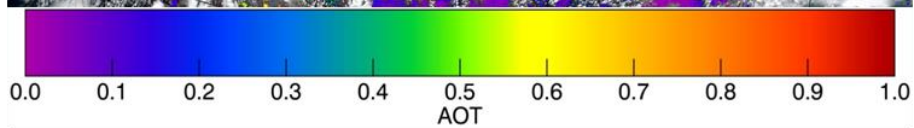
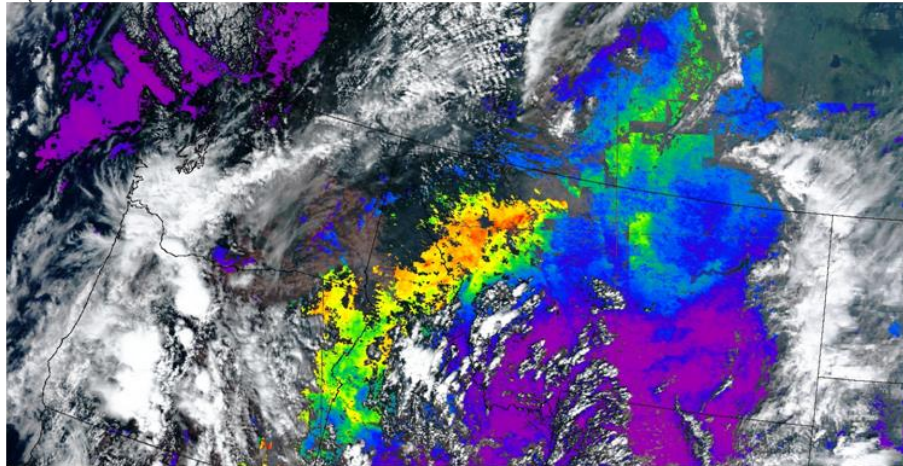
(a) VIIRS True-color Aug. 25<sup>th</sup>, 2013



(b) VIIRS IP AOT 20:34 - 20:37 UTC



(c) MAIAC AOT 20:20 - 20:25 UTC



763

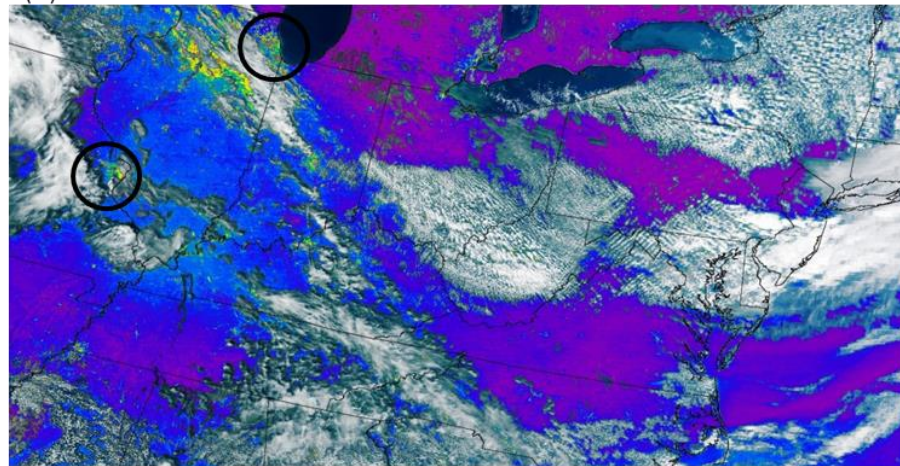
764 Figure 9. Example of high aerosol loading on August 25<sup>th</sup>, 2013 over the western U.S. due to regional  
765 fires. (a) True-color image from S-NPP VIIRS; (b) VIIRS high quality IP AOT; (c) MAIAC AOT.



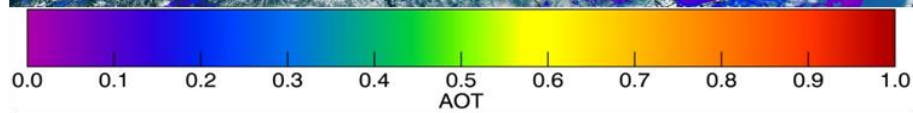
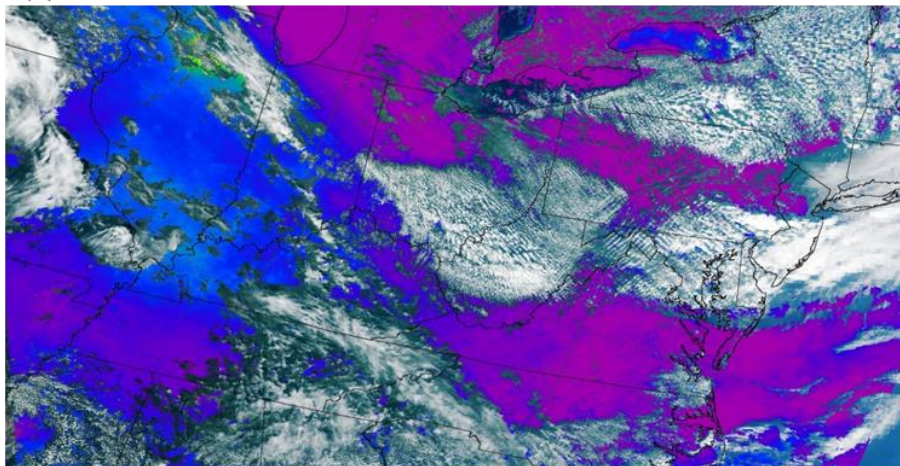
(a) VIIRS True-color Sept. 5<sup>th</sup>, 2013



(b) VIIRS IP AOT 18:43 – 18:45 UTC



(c) MAIAC AOT 18:20 – 18:25 UTC



766

767 Figure 10. Image of a moderate AOT case from September 5, 2013. (a) True-color image from S-NPP  
768 VIIRS; (b) VIIRS high quality IP AOT; (c) MAIAC AOT.